## The Myopia of Imperfect Climate Models: The Case of UKCP09

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The United Kingdom Climate Impacts Program's UKCP09 project makes high-resolution forecasts of climate during the twenty-first century using state of the art global climate models. The aim of this article is to introduce and analyze the methodology used and then urge some caution. Given the acknowledged systematic errors in all current climate models, treating model outputs as decision-relevant probabilistic forecasts can be seriously misleading. This casts doubt on our ability, today, to make trustworthy, high-resolution predictions out to the end of this century.

**1. Introduction.** There is now a widespread consensus that global warming is real and in large part due to human activities.<sup>1</sup> But knowing that the climate is getting warmer is of limited use in designing detailed adaptation strategies.<sup>2</sup> The impact of climate change on humans occurs at a local scale, and so ideally we would like to know what changes we have to expect in our

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1. The existence of a widespread consensus is documented in Oreskes (2007); the evidence for the warming being anthropogenic is documented in the last IPCC Report. Throughout, "IPCC" refers to Solomon, Qin, and Manning (2007).

2. It may well be enough for mitigation: knowing roughly what is likely to happen may be reason enough not to go there.

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immediate environment; reliable answers would greatly aid decision makers (Oreskes et al. 2010; Smith and Stern 2011; Sexton et al. 2012; Tang and Dessai 2012).

The United Kingdom Climate Impacts Program's UKCP09 project aims to answer exactly such questions by making high-resolution forecasts of twenty-first-century climate.<sup>3</sup> It generates its predictions using state-of-theart global climate models. The Intergovernmental Panel on Climate Change (IPCC) has confidence that these models have some skill at continental scales and above. This leaves open the question whether decision-relevant high-resolution predictions can be constructed with today's models.

The aim of this article is to introduce and analyze the methodology used by UKCP09 and then urge some caution. Given the acknowledged systematic errors in all current climate models, treating model outputs as the basis for decision-relevant probabilistic forecasts can be seriously misleading (Stainforth et al. 2007). This casts doubt on our ability, today, to make trustworthy, high-resolution predictions out to the end of this century.<sup>4</sup> Here we introduce the aims of UKCP09 (sec. 2), outline the method used to generate predictions (sec. 3), discuss the project's handling of structural model error (sec. 4), argue that crucial assumptions are untenable (sec. 5), and then draw some conclusions (sec. 6).

**2.** UKCP: Aims and Results. The declared aim of UKCP09 is to provide decision-relevant forecasts on which industry and policy makers can base their future plans:

To adapt effectively, planners and decision-makers need as much good information as possible on how climate will evolve, and supplying this is the aim of . . . UKCP09. They are one part of a UK government programme of work to put in place a new statutory framework on, and provide practical support for, adaptation.

3. "UKCP09" stands for "United Kingdom Climate Projections 2009," where "2009" indicates that it was launched for public use in 2009. UKCP09 is documented in the Briefing Report (Jenkins et al. 2009), the Science Report (Murphy et al. 2010), and two recent papers, Sexton et al. (2012) and Sexton and Murphy (2012). The full set of predictions is at http://ukclimateprojections.defra.gov.uk/.

4. In this article we use the word 'trustworthy' to denote probability forecasts that one might rationally employ for decision-making purposes using probability theory in the standard way. Such probability forecasts are expected to be robust and reliable, the kind a good Bayesian would make. We wish to avoid the kind of analysis that inspired Rubin's remark that "a good Bayesian does better than a non-Bayesian, but a bad Bayesian gets clobbered" (cited in Good 2009, 139). There may be many justifiable and interesting scientific reasons to construct probability forecasts; our criticism of them in this article is only in regard to their direct use in decision support (as, e.g., illustrated in the worked examples of UKCP09).

The projections have been designed as input to the difficult choices that planners and other decision-makers will need to make, in sectors such as transport, healthcare, water-resources and coastal defences, to ensure the UK is adapting well to the changes in climate that have already begun and are likely to grow in future. (Jenkins et al. 2009, 9)

In a system as complex as the world's climate, it is absurd to produce point forecasts (i.e., forecasts saying that a particular event will happen at a particular time with certainty). UKCP09 produces what they dub Bayesian probability forecasts, which "assign a probability to different possible climate change outcomes recognizing that . . . giving a range of possible climate change outcomes is better, and can help with making robust adaptation decisions, but would be of limited use if we could not say which outcomes are more or less likely than others" (Jenkins et al. 2009, 23).

The challenges many decision makers have to address arise at a local level: flood barriers have to be built in a particular location and to a given height, and so on. For this reason, local user-relevant information about the impacts of climate change is the most useful, assuming of course that it is not misinformative (Smith and Stern 2011).

UKCP09 strives to meet the demand for decision-relevant information at the local level by producing highly specific information (Jenkins et al. 2009, 6–7). Probabilities are given for events on a 25-kilometer grid. Forecasts are made for finely defined specific events such as changes in the temperature of the warmest day of a summer and the precipitation of the wettest day of a winter. It is predicted, for instance, that under a medium emission scenario the probability for a 20%–30% reduction in summer mean precipitation in central London in 2080 is 0.5 (Jenkins et al. 2009, 36).

**3.** The Architecture of UKCP09. These predictions are generated with a method involving both global climate models (GCMs) and elaborate statistical techniques. In this section we outline the method, in five parts, with the aim of identifying key assumptions and making its architecture visible.<sup>5</sup>

3.1. Modeling. The cornerstone of UKCP09 is HadCM3, a GCM developed at the Hadley Centre. The model consists of two coupled modules, one representing the earth's atmosphere (including land surface processes and surface-atmosphere exchanges) and one representing the oceans. Our best descriptions of these fluids come from nonlinear partial differential equations (PDEs), which define the evolution of continuous fields representing the atmosphere or ocean. It is neither possible to integrate PDEs exactly nor to measure precisely the continuous fields required to initialize them. Instead

5. Our account of the method is based on Murphy et al. (2010, chap. 3).

they are discretized onto a grid in space and in time. Today's computational constraints force climate models to use a relatively coarse grid; those used in UKCP09 have a typical resolution of around 300 kilometers.

The model includes tens of thousands of dynamical variables and hundreds of parameters specifying the value of physical magnitudes, representing physical constants or controlling small-scale processes that are not resolved explicitly. To aid the discussion let us introduce some notation. Let  $x(t) = \{x_1(t), x_2(t), \ldots\}$  be the vector of all dynamical variables and  $\alpha = \{\alpha_1, \alpha_2, \ldots\}$  the vector of all parameters in the model; then let  $\phi_t^C(x; \alpha)$  be time evolution of HadCM3, specifying the future value of the system's dynamical variables given certain initial conditions and certain parameter values.

Even state-of-the-art computers take a long time to make a 100-year run of  $\phi_t^C(x; \alpha)$ , and so a simpler model is needed for most calculations (a 'run' is the calculation of the future value of *x* given a particular initial condition and a set of parameter values). To this end the entire ocean module is eliminated and replaced by a so-called slab ocean. With no currents and a uniform effective depth, this slab ocean is defined via simpler equations. The result is HadSM3.<sup>6</sup> We write  $\phi_t^S(x; \alpha)$  to denote the time evolution of this model, where we take it as understood that the vectors *x* and  $\alpha$  vary with the model structure (HadSM3 having fewer variables and parameters than HadCM3).

3.2. Perturbed Physics Ensemble. The problem in determining the future values of x is that "the available information is seldom precise enough to allow the appropriate value of a given parameter to be accurately known" (Murphy et al. 2010, 37). Not knowing what value of  $\alpha$  to use in our calculations, assuming there is one, "gives rise to the parameter component of model error" (37).<sup>7</sup>

The technique of a *perturbed physics ensemble* (PPE) is designed to address this difficulty.<sup>8</sup> The idea of a PPE is to calculate future values of x for a number of different values of  $\alpha$ . If, for instance, modelers are uncertain about "the" reasonable value of parameter  $\alpha_2$  but believe that it lies between  $a_{2,\min}$  and  $a_{2,\max}$ , they carry out calculations of x for as many values in the interval  $[a_{2,\min}, a_{2,\max}]$  as they can afford. The variability of the outcomes then gives them a sense of the sensitivity of the model. Calculating future x's for a number of different parameter values amounts to construct-

6. Going from HadCM3 to HadSM3 roughly doubles the speed of the model.

7. This assumption is controversial. Smith (2002, 2006) argues that for imperfect models appropriate values (leading to trustworthy forecasts) may not exist. For want of space we set these worries aside; for more on this point, see Smith and Stern (2011).

8. We note in passing the lack of unanimity on whether the second 'P' of PPE stands for 'parameter', 'parameterization', or 'physics'. For more on PPEs, see Allen and Stainforth (2002).

ing a PPE because the variation of the parameter values amounts to perturbing the physics yet without changing the mathematical structure of the model (because the functional form of each equation remains unchanged).

In a complex model like HadCM3 a single research center can only make a relatively small number of runs due to the limitation of computational resources. The question then is how to construct a PPE for a model with hundreds of parameters if only a small number of runs can be made. UKCP09 solves this problem by first restricting attention to atmospheric parameters and then soliciting parameterization experts to identify those parameters that control the crucial processes in the system and on which the future values of x depend most sensitively. This process led to the identification of 31 crucial parameters and the definition of associated plausible intervals for them. To explore the uncertainty of future values of x brought about by the variation in these 31 parameters, 280 runs were made with HadSM3. Information from 17 HadCM3 runs was added later.

3.3. *Emulator*: Unfortunately this number remains too small to provide a good understanding of the diversity of outcomes. An *emulator* is therefore built to provide values of x corresponding to values of  $\alpha$  for which no runs were made. In other words, the emulator "fills the gaps" between the 280 points obtained in section 3.2.

3.4. Probability. What is the uncertainty of future values of x given the diversity in  $\alpha$ ? Uncertainty is quantified by giving a probability distribution over the interval associated with  $\alpha$ . The emulator correlates every value of  $\alpha$  with an outcome x, and these distributions are reinterpreted as probabilities for x. UKCP09 assumes that each value of  $\alpha$  within the middle 75% of the interval is equally likely and that the probability linearly drops to zero at the minimum and maximum values. These probabilities are then adjusted by assigning relative weights to all values of  $\alpha$  according to the emulator-implied ability of the model to represent observations when simulating a period similar to the past.

*3.5. Downscaling.* The model calculations are done with a resolution of the order of 300 kilometers but predictions are sought at the 25-kilometer scale. To generate predictions at that level of detail the results are down-scaled using simulations of a limited area regional climate model configured from HadCM3 and run at a 25-kilometer horizontal resolution.

The endeavors of these five parts taken together produce the predictions we have discussed in the last section.

**4. Structural Model Error.** Each of these steps raises potentially significant conceptual and methodological questions. For want of space we can

only deal with what we see as one central problem in section 3.1: structural model error (SME). Like every model, HadCM3 has its imperfections. In order to specify  $\phi_t^C(x; \alpha)$  a number of strong idealizing assumptions are made. These include (effective) distortion of the topography of the earth (mountain ranges like the Andes are systematically too smooth and too short; small volcanic island chains with visible impacts on atmospheric circulation do not exist) and approximations of the effects of cloud fields that cannot be simulated realistically at the available resolution. Furthermore, solutions of the discretized PDE differ from those of the original PDE, and the PDE itself differs from what the true equations of the world would be (assuming such equations exist at all). In addition, there are limitations to our scientific understanding of the climate system, and there may be relevant factors and processes about which we are simply unaware that would lead us to alter the equations of the model even under our current computational constraints.

Inasmuch as SME is due to shortcomings in the equations of the model, the challenges it poses to forecasting cannot be resolved by varying the model's parameters. If a model has SME, this means that the time evolution of an ensemble will, eventually, differ from that of a better model and indeed from reality itself, if a relevant distribution can be associated with reality. No adjustment of the parameters can remove this difference. One crucial question is: how soon do dramatic effects of SME manifest themselves in a given situation? Another is what extent can a model with SME still be informative about the target system? On what timescales does the science (which underlies the model) suggest that a decision maker should expect a big surprise if he took the model outputs as trustworthy?

UKCP09 acknowledges the presence of SME and proposes a way to deal with it. The message is that the uncertainties due to SME can be estimated and taken into account in projections.<sup>9</sup> In this section we outline their approach, and in the next we ask whether its use for the provision of quantitative decision support is justified.

UKCP09 aims to capture the difference between the model and the real world with a so-called discrepancy term, which "represents how informative the climate model is about the true climate, and it measures the difference between the climate model and the real climate that cannot be resolved by varying the model parameters. Such differences could arise from processes which are entirely missing from the climate model, or from fundamental deficiencies in the representation of processes which are included, through (say) limited resolution or the adoption of an erroneous assumption in the parameterisation scheme" (Sexton et al. 2012, 2515; see also Murphy et al. 2010, 63–64). Assume, then, that we are interested in 'the true climate' at a

9. The UKCP09 science report calls the proposed method "an appropriate means of quantifying uncertainties in projected future changes" (Murphy et al. 2010, 66).

particular future instant of time  $t^*$  (e.g., August 2080),<sup>10</sup> and let *c* be the true value of *x* at  $t^*$  (hereafter 'target'). The relation between the model output and the target then is:<sup>11</sup>

$$c = \phi_{t*}^{\mathcal{S}}(x_0; \alpha^*) + d,$$

where  $\alpha^*$  is the set of parameter values that best simulates the target. The discrepancy *d* is a vector in the system's state space, and it can be interpreted as telling us "what the model output would be if all the inadequacies in the climate model were removed, without prior knowledge of the observed outcome" (Sexton et al. 2012, 2515).

Now follow two crucial assumptions. The first assumption is "that the climate model *is* informative about the real system and the discrepancy term can be seen as a measure of how informative our climate model is about the real world" (Sexton et al. 2012, 2515; original emphasis). The idea is that informativeness comes in degrees and is indirectly proportional to the length of *d*: the smaller *d* the more informative the model. That the model is informative then amounts to assuming that *d* is small. We call this the *informativeness assumption*. The second assumption concerns the discrepancy. While *d* is defined as a vector, in practice one cannot know the exact vector and so it is assumed that there is probability distribution  $\varepsilon$  over *d*. This distribution is then assumed to be Gaussian (Sexton et al. 2012, 2515). We refer to the package of the two as the *core assumption*.

With this assumption in place UKCP09 sets out to estimate the parameters of  $\varepsilon$ . Not being omniscient, one cannot just compare model outputs with the truth. The crucial move in UKCP09 is to use a multimodel ensemble (MME) as a proxy for the truth: "Our key assumption is that sampling the effects of structural differences between the model chosen for the PPE and alternative models provides a reasonable proxy for the effects of structural errors in the chosen model relative to the real world" (Sexton et al. 2012, 2516).<sup>12</sup>

The MME in question contains 12 models (Sexton et al. 2012, 2519). The claim then is that measuring the average distance of HadSM3 to a set of different models yields a similar result as measuring its distance to the real world—hence, d can be determined by measuring by how much HadSM3

11. See Sexton et al. (2012, 2521). Throughout we use our own notation, which differs from Sexton's.

12. See also Sexton et al. (2012, 2526) and Murphy et al. (2010, 64).

<sup>10.</sup> There is, obviously, a serious confusion that we cannot clarify here: climate is a distribution, the state of the atmosphere at a given time ('weather') is a point. Our uncertainty in the 'true' (we would prefer 'target') value of x at time t does not correspond to the climate at time t.

diverges from those other models. We call the view that an MME is a trustworthy proxy for the real world the *proxy assumption*.

For each model in the ensemble  $\alpha^*$ , the best HadSM3 analogue, is determined. Having found the best analogue, the prediction error *b* is calculated, essentially the difference between the two model outputs. With these *b*'s the mean and variance of  $\varepsilon$  are determined.<sup>13</sup>

Under the proxy assumption, this procedure quantifies the additional uncertainty due to model error. This uncertainty is now added to the uncertainty about values of x obtained in section 3.4, yielding the total uncertainty. The uncertainty is expressed as a probability distribution; it is this distribution that is presented as guidance for decision makers.

In the following section we argue that neither the core assumption nor the proxy assumption are well founded. The restriction to these two assumptions is due to lack of space; other aspects of the approach also raise serious questions.

5. The Assumptions Scrutinized. Our discussion of the core assumption focuses on informativeness. It is an undisputed fact that systematic errors in the models in question lead to nontrivial macroscopic errors of simulation, of the past and of the future. Seager et al. (2008) have noted their inability to reproduce the dust bowl of the 1930s even given the observed sea-surface temperatures. This is not a small inadequacy when one is focused on the resolution offered by UKCP09. Given these systematic errors, there are lead times at which the failure of the model to simulate realistic weather causes the climate of the model to differ from that of the planet (Smith and Stern 2011). Inasmuch as the models used are not close to the target, the informativeness assumption fails. Figure 1 shows model global mean temperatures over the last century of the 24 models of the Coupled Model Intercomparison Project Phase 3 (CMIP3).<sup>14</sup> Note that while all models show warming between 1900 and 2000, their average temperatures vary tremendously. The magnitude of the error in the global mean in a hindcast of the last century casts significant doubt on the viability of the informativeness assumption on a 25-kilometer forecast to the end of this century.

Even if one were to discard the above as undue pessimism and uphold the informativeness assumption, there is a further problem. An argument to uphold the informativeness assumption *now* must be based on the trustworthiness of the modeling assumptions, typically taking the form that if the model assumptions are close to the truth, then the model outputs must be close to the truth too. While this inference works in some specific applica-

13. See Sexton et al. (2012, 2521–27). To be precise, what is determined is the co-variance matrix.

14. Thanks to Ana Lopez for producing the figure.

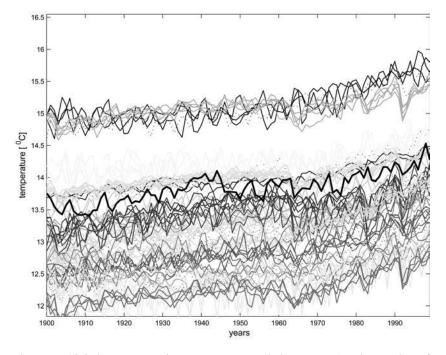


Figure 1. Global mean annual temperature, twentieth century. A color version of this figure is available online.

tions, it is generally false for nonlinear models even if their SME is extremely small (Frigg et al. 2014). Since the relevant climate models are nonlinear, it follows that even if the model assumptions were close to the truth this would not automatically warrant trust in the model outputs. In fact, the outputs for relevant lead times 50 years from now could still be seriously misleading.

The discussion of the proxy assumption is complicated by the fact that the literature on the subject exhibits a certain degree of schizophrenia. On the one hand, the method is illustrated and advertised as delivering trust-worthy results; on the other hand, disclaimers that effectively undermine the crucial assumptions are also included, sometimes parenthetically, obscurely, or deep within technical discussions.<sup>15</sup> The documentation gives with one hand and takes back with the other. We now review the activities of both hands and conclude that the hand that takes back voids the trustworthiness of the forecasts for quantitative decision support.

The first reason cited in support of the proxy assumption is that multimodel averages give a better representation of climate than any individual

15. An example is Murphy et al. (2010, 63-69).

model: "Indeed, the multimodel ensemble mean has been shown to be a more skillful representation of the present-day climate than any individual member" (Sexton et al. 2012, 2526). Yet it is also acknowledged that "systematic errors to all current climate models persist" (2526) so even if one were to accept that such a multimodel mean were more skillful at representing the present day climate than individual models, is 'more skillful' close to being 'skillful'? Unfortunately, there appears to be no evidence that 'more skillful' can be equated with 'skillful' for many variables of importance for future climate change.

The second reason mentioned in support of the proxy assumption is that "the structural errors in different models can be taken to be independent" (Murphy et al. 2010, 66) and that therefore the ensemble samples uncertainty well. However, immediately after we are warned that "whilst there is evidence for a degree of independence . . . , there is also evidence that some errors are common to all models . . . , due to shared limitations such as insufficient resolution or the widespread adoption of an imperfect parameterisation scheme. From this perspective, our estimates of discrepancy can be viewed as a likely lower bound to the true level of uncertainty associated with structural model errors" (Murphy et al. 2010, 66). And then the conclusion is drawn that "the main (and inevitable) limitation, however, is that it [the proxy assumption] does not account for the potential impacts of errors common to all climate models used in the prediction" (Sexton et al. 2012, 2516).

One cannot have one's cake and eat it too. If there are common errors, the proxy assumption fails. Indeed such common errors have been widely acknowledged (see, e.g., Parker 2011). Furthermore, the mathematical space of all possible climate models (if there is some such thing) is huge, and there is no reason to believe that the 12 models we de facto work with provide a representative sample.

For these reasons, the assumption that the use of an MME will accurately quantify the distance to our true target is unjustified. It produces a distribution that is more consistent with the diversity of current models but need not reflect the uncertainty in our future (Smith 2002). It is important to note that the fear is not so much that the width of the uncertainty distribution is too narrow but, rather, that the distribution is simply in the wrong place and that the mean of the distribution will shift significantly if the model simulations become realistic.

Echoing Murphy and colleagues, we note that "it is important to stress that our approach to the specification of discrepancy can only be expected to capture a subset of possible structural modelling errors and should be regarded as a lower bound" (Murphy et al. 2007, 2011). A lower bound need neither yield trustworthy forecasts nor provide a suitable basis for quantitative decision support. **6.** Conclusion. We have argued that there is little evidence for interpreting UKCP09's predictions as trustworthy forecasts for quantitative decision support. Questioning the evidence, however, does not amount to proving it wrong. Our point is that the premises of the argument do not warrant trust in the results, and for decision support in the face of climate change this is the crucial aspect.

To be fair to the scientists who worked very hard to make UKCP09 the best it could be, several points should be noted. First the deliverables of the project were defined before any viable approach to meet them was available in the peer-reviewed literature. Second, the United Kingdom Climate Impacts Program, which is much broader than UKCP09, faced the dilemma of motivating users to engage with the real challenges and significant risks posed by climate change in the face of deep uncertainty: the challenge of keeping users interested when the information they most desire lies beyond the reach of today's science. And last, pointers to the fact that a naïve interpretation of UKCP09 probability distributions is untenable can indeed be found within the UKCP09 material.

That said, the aim of UKCP09 was to provide trustworthy forecasts now, and this, we have argued, they fail to do. Kelly's (1979) plea holds today: climate prediction experiments remain essential, and they must "be conducted with scientific rigour and presented with an honest assessment of the uncertainties involved" (Kelly 1979, 182).

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